



# Nonparametric analysis of extreme precipitation with long return periods in Japan – using d4PDF

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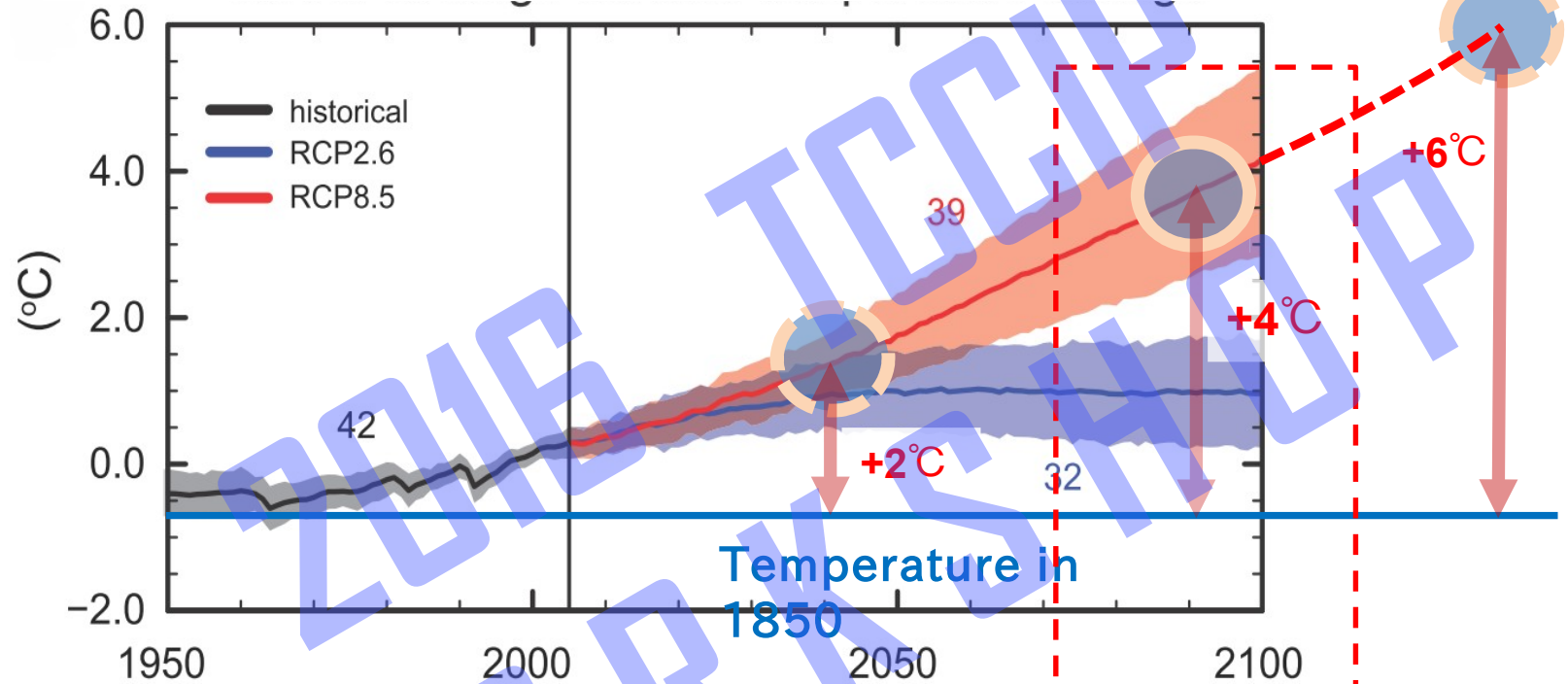
2 Meteorological Research Institute, Japan

# Introduction

- Estimation of extreme precipitation is a important task to avoid disasters like floods, land-slide, and so on.
- This study focus on rare events such that they occur once, on average, in a given years (called **return period**)
- Analyses of extreme precipitation are usually based on **annual maximum of daily precipitation** (called **R1d** hereafter).
- However, observation or previous climate experiments yield a few samples of R1d (the order of tens or hundreds).
- Using d4PDF dataset, which has the data of the order of thousands, we improve the reliability of the estimate of extreme precipitation and reveal their **statistical features**.

# Design of d4PDF experiments

Surface temperature of the whole world mean



60km AGCM

100 members

90 members  
( $6\Delta T \times 15\delta T$ )

NHRCM 20km  
(Japan Islands)

50 members

3000 years  
samples

90 members

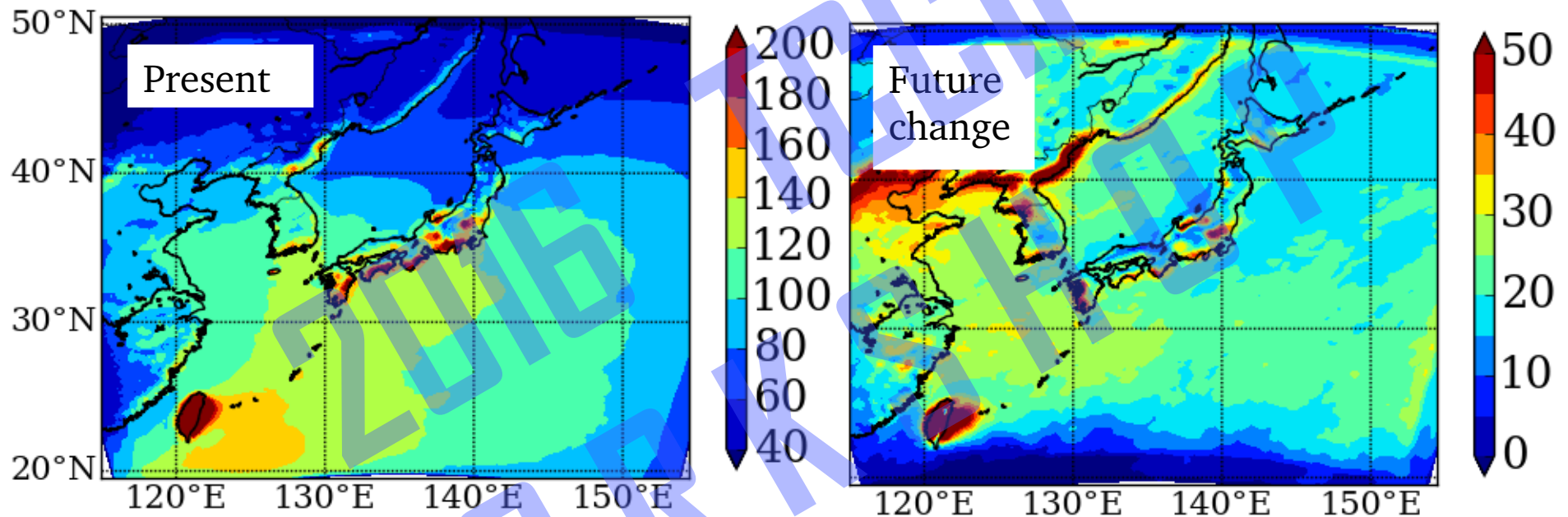
5400 years  
samples

1951 Present exp. 2010

60 years

# Extreme precipitation (once a year - R1d) from regional model of d4PDF

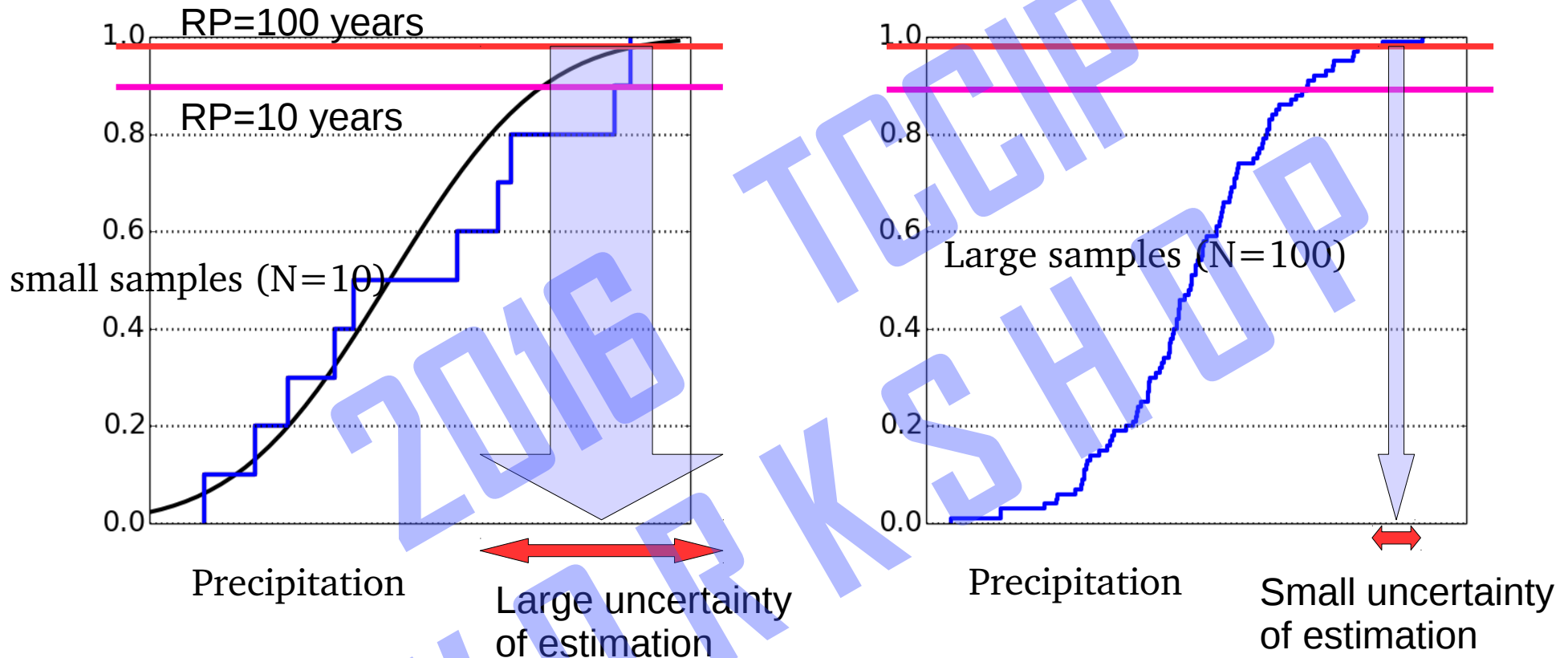
Ensemble mean of present R1d (mm) and future change (mm)



- Although R1d tends to decrease along the latitude in this domain, large values of R1d emerges on the land regions, particularly in Taiwan and on the coast of Pacific ocean of Japan islands.
- Geographic distribution is well reproduced thank to the downscaling to the regional model.

# Return levels of extreme precipitation

Empirical cumulative distribution function (CDF) of R1d



- The estimation of extreme precipitation is difficult with small sample numbers.
- To solve this difficulty, we have two choices:
  - (A) well-known functions (e.g., GEV) are fitted to empirical CDF (assumption about the characteristics of population of R1d data)
  - (B) increase the number of samples.

# Two methods to calculate a return level when given a return period

## (A) GEV (generalized extreme value) method

- Fit a GEV function to empirical CDF and read the precipitation value of the intersection point of the GEV function and the cumulative probability line which is determined by given return period.
- Fitting of the GEV distribution is performed using the maximum likelihood method.
- The uncertainty of the estimate of the return levels is derived from the uncertainty of the GEV fitting.

## (B) Non-parametric method

- Read the precipitation value of intersection point of the empirical CDF and the cumulative probability line directly.
- The uncertainty of the estimate is obtained with the bootstrap method (one of sub-sampling methods)

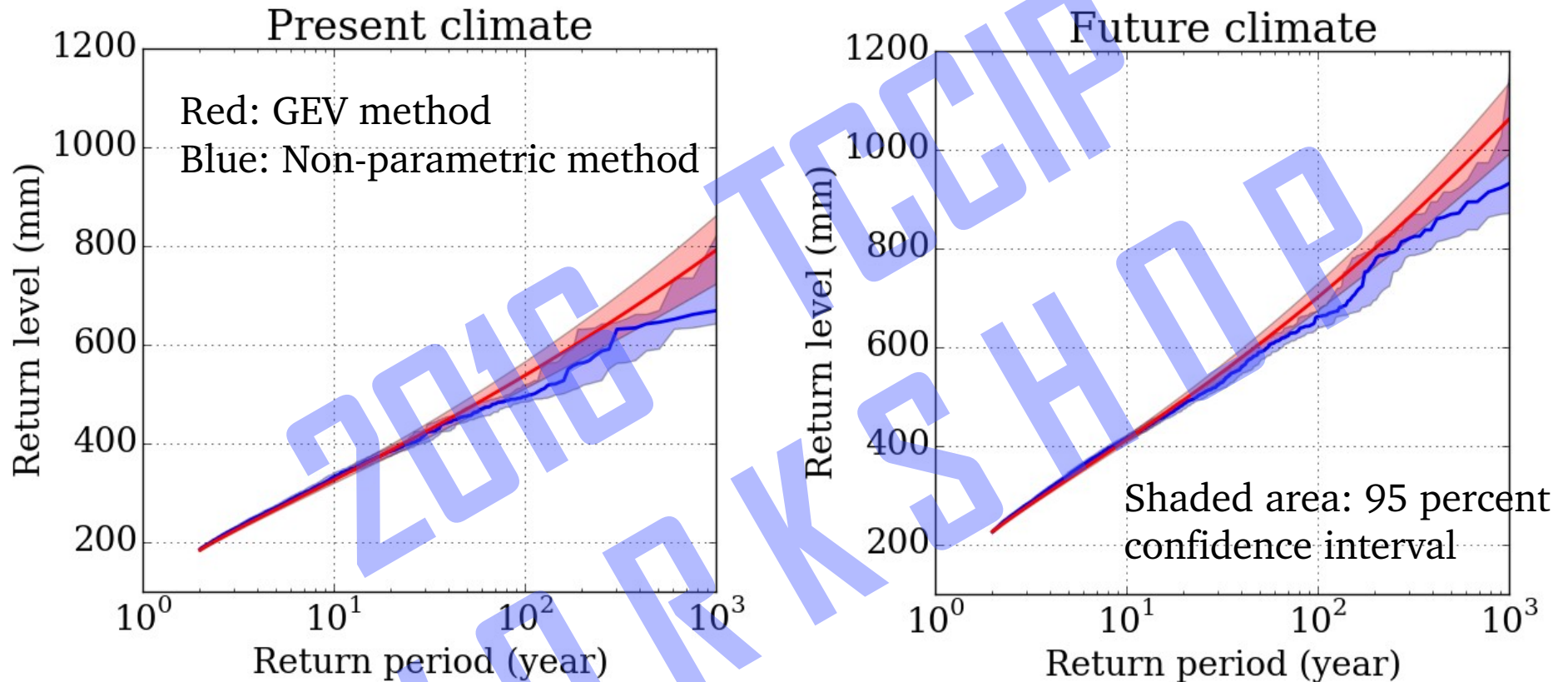
Estimation from the non-parametric method is more reliable than that from the GEV method because no assumption about the statistical features of the R1d is made.

2018 TCCIP  
WORKSHOP

# Results



# Results of return levels of precipitation (at Taipei for example)

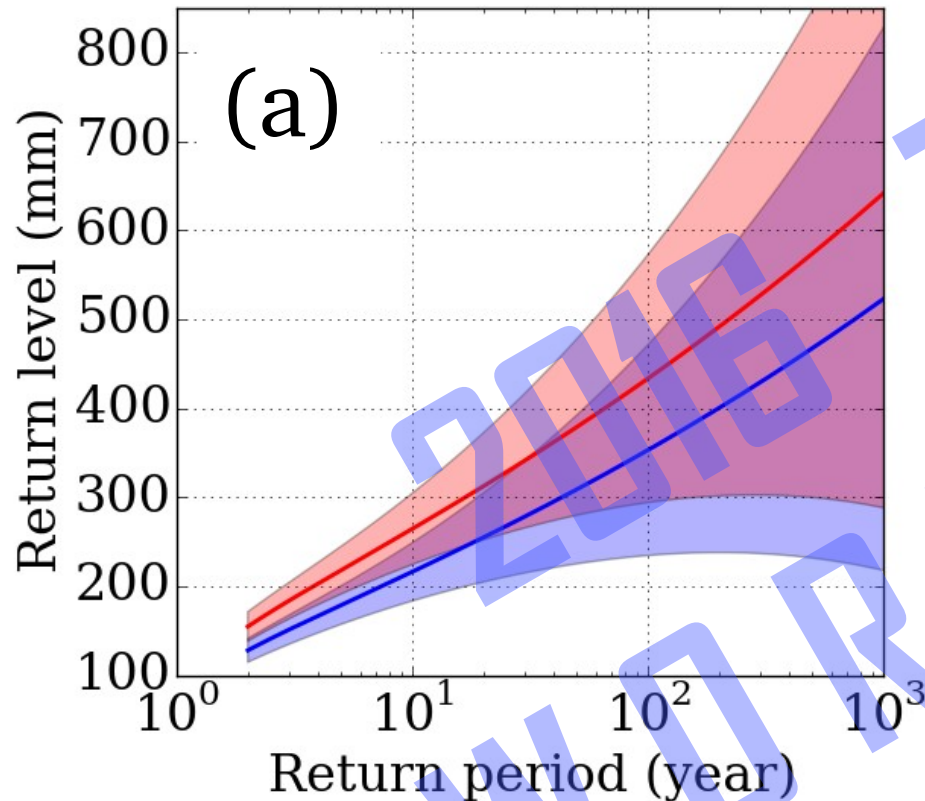


- GEV method and non-parametric methods yield approximately the same results about the return levels, which implies that the assumption of GEV fitting is reasonable for Taipei R1d data.
- Both methods project a increase of the extreme precipitation.

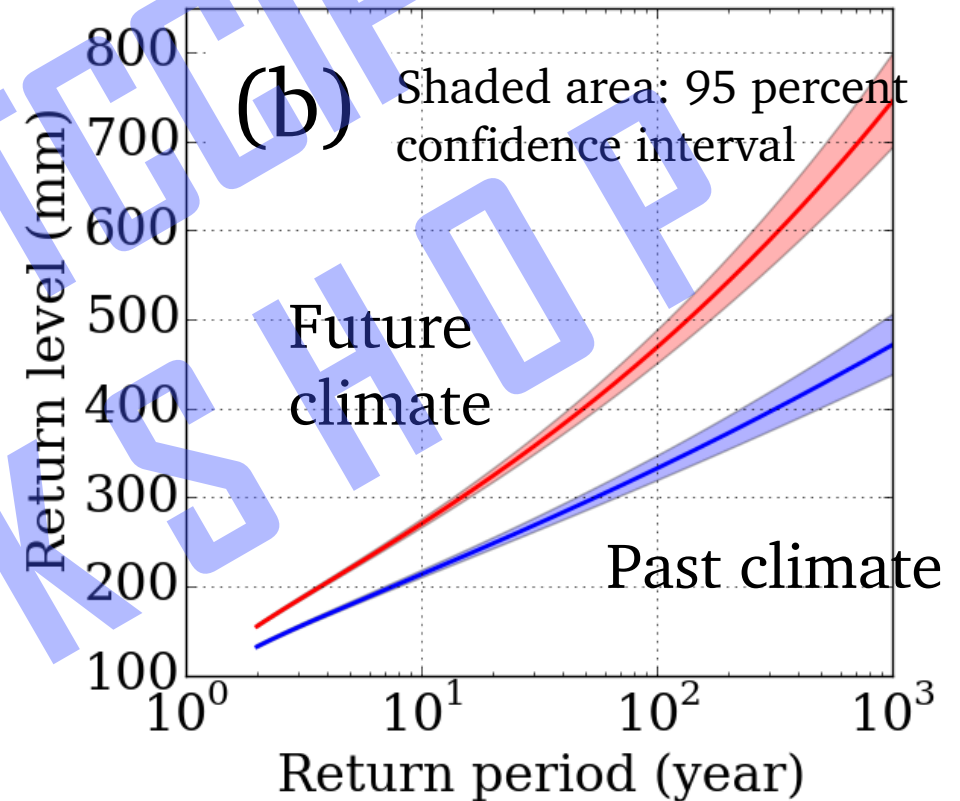


# Benefits of large ensemble members from d4PDF at Tokyo as an example

One experiment  
(61 years samples)



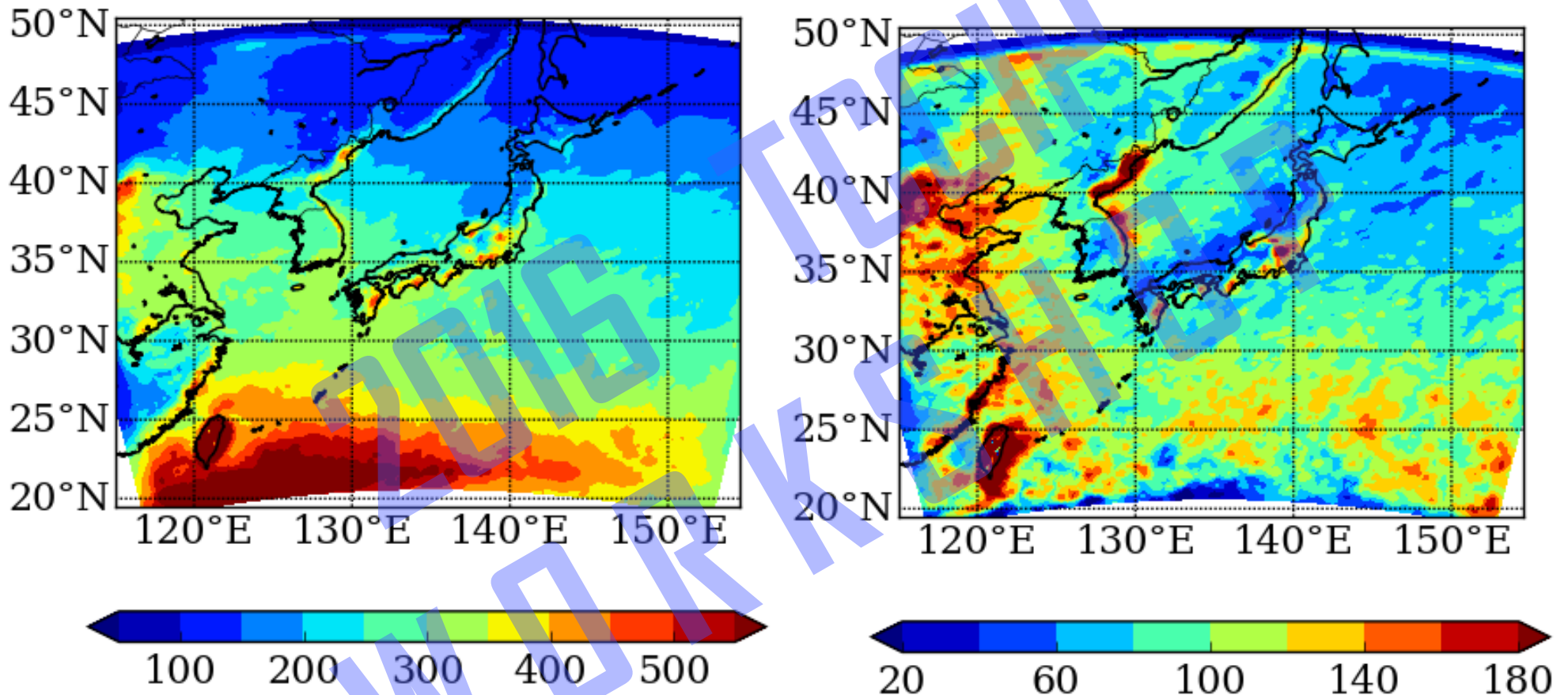
All ensemble members of d4PDF  
(Past: 3050 years, Future: 5490years)



Confidence intervals of return levels shrink thanks to the large samples from d4PDF, leading to the reliability of estimation and enabling the detection of future changes of extreme precipitation in a statistical sense.

# Geographical distribution of extreme precipitation

(left) Present return levels (mm) with 100 years return periods and  
(right) their future changes



- The geographical pattern of the 100 years return levels is similar to that of R1d.
- However, the value of precipitation is much larger.
- The dynamical downscaling enables the analyses of extreme precipitation focusing on the detailed geographical distribution e.g., orographic precipitation.

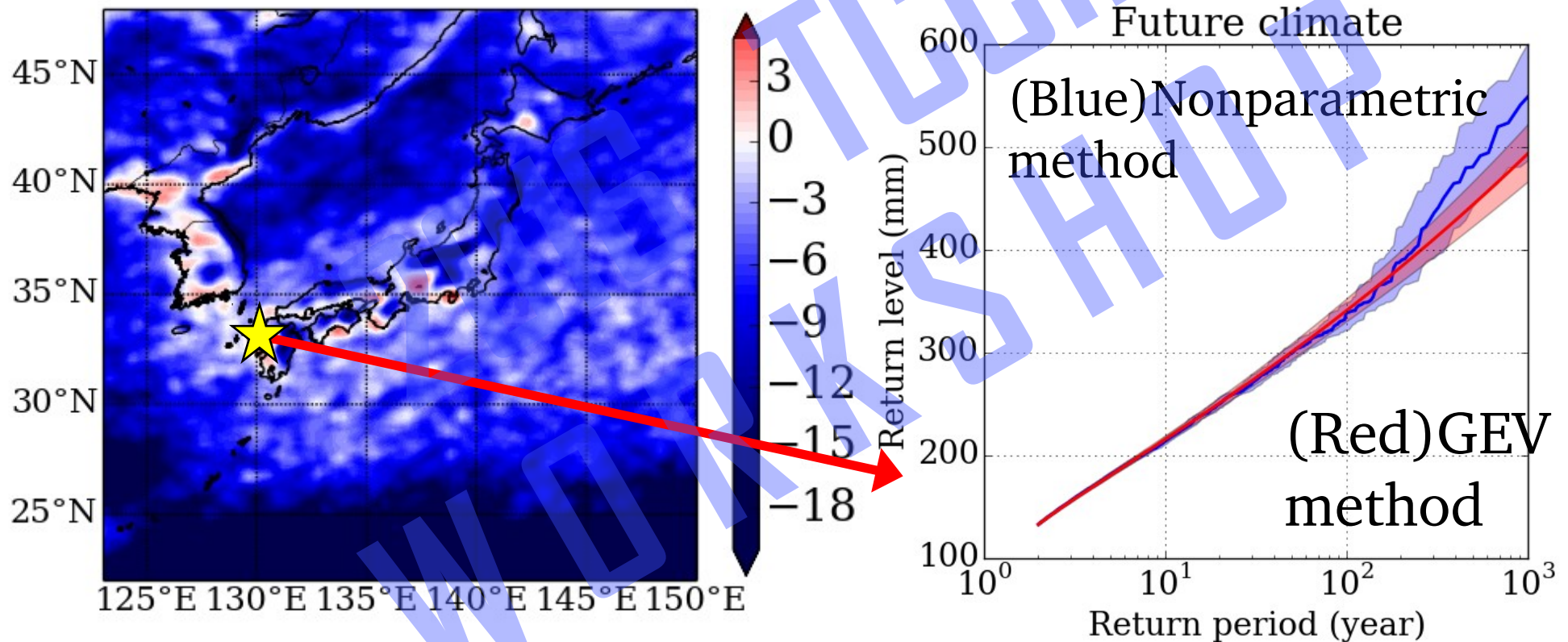
# Validity of the GEV assumption

2016 TCCIP  
WORKSHOP

# Comparison between the results of two methods: GEV method and nonparametric method

(Left) Relative difference (%) of 100 years return levels between the results by GEV method and nonparametric method.

Red area indicates under-estimate of GEV method and blue area is over-estimate.



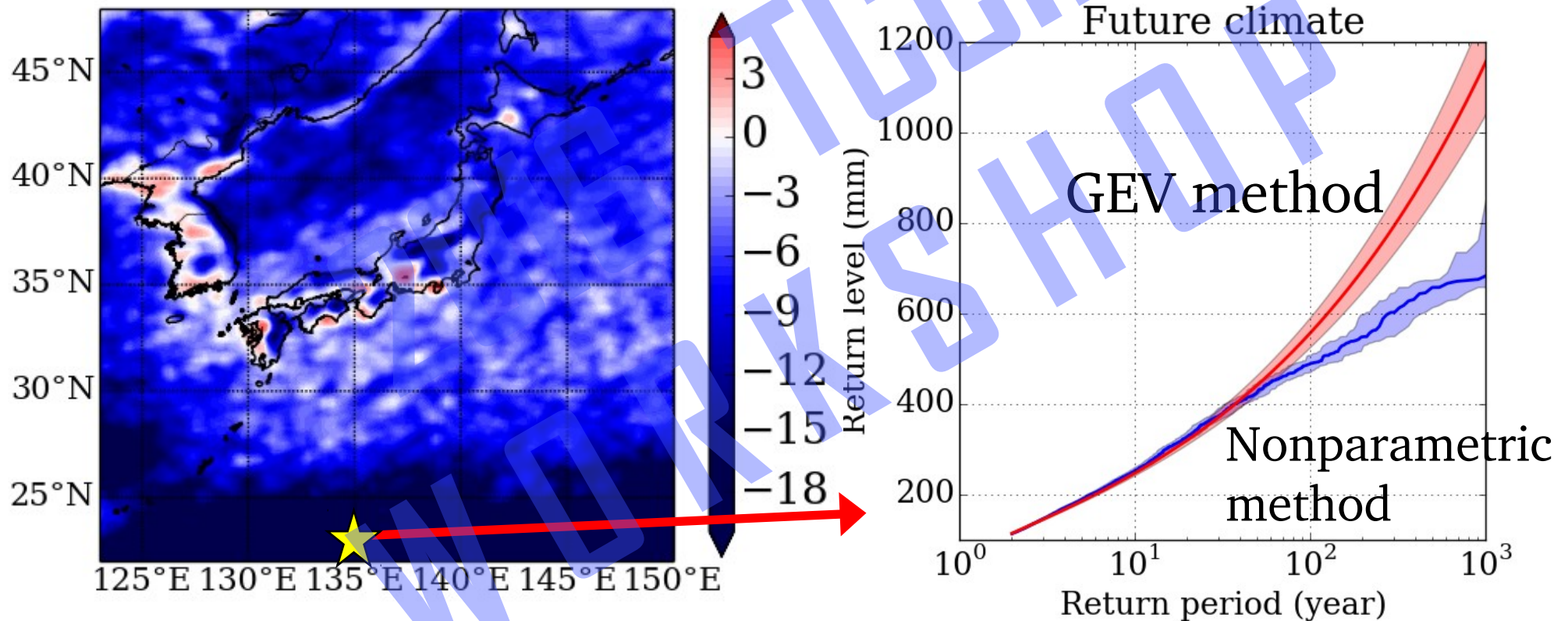
- The reliability of GEV method is relatively good around Japan islands and Korean peninsula.
- The GEV method tends to yield overestimated return levels on the ocean.



# Comparison between the results of two methods: GEV method and nonparametric method

(Left) Relative difference (%) of 100 years return levels between the results by GEV method and nonparametric method.

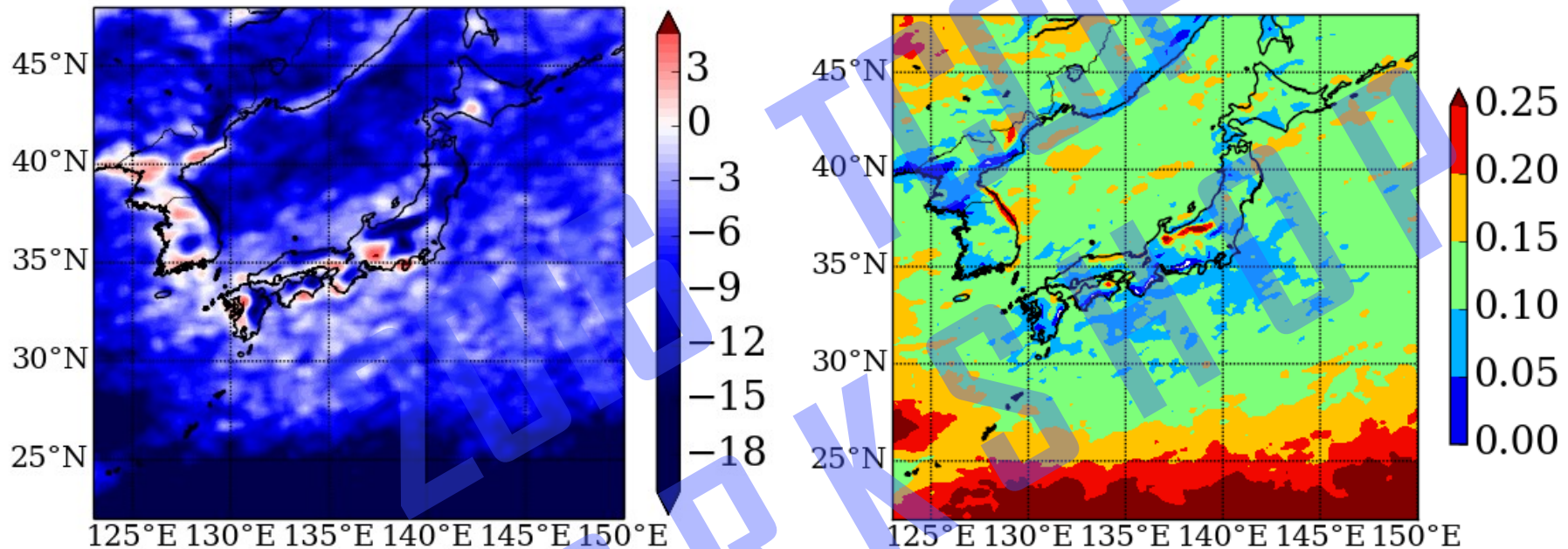
Red area indicates overestimation of GEV method and blue areas are underestimation.



- The reliability of GEV method is relatively good around Japan islands and Korean peninsula.
- The GEV method tends to yield overestimated return levels on the ocean.

# Why the fitness of GEV to R1d is low?

Shape parameter  $\xi$  of GEV (right figure); large  $\xi$  means heavy tailed GEV distribution.



The areas where overestimation of return levels are observed are overlapped with those of high shape parameters.

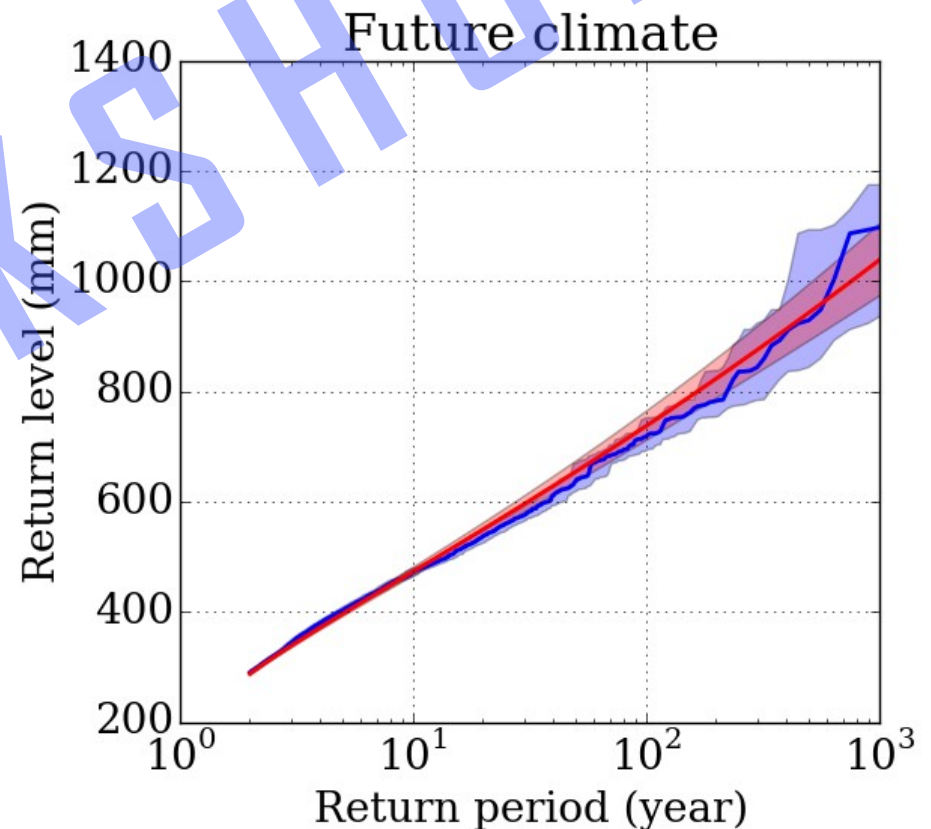
- (i) Bad fitting of GEV to R1d distribution
- (ii) Artificial large shape parameter and heavier tail of GEV.
- (iii) Overestimation of return levels with large return periods.

# Why the fitness of GEV to R1d is low? (cont.)

- R1d is not a perfect extreme value.  
The assumption that a value is followed by GEV is that this value is the maximum value of large enough samples.
- This is considered as follows, on land, orographic extreme precipitation invariably happens at least once in a year, while on ocean, extreme precipitation does not occur without a direct hit of a tropical cyclone

GEV fitting test on the southern ocean  
R10d : maximum daily precipitation  
over 10 years, not 1 year (R1d)

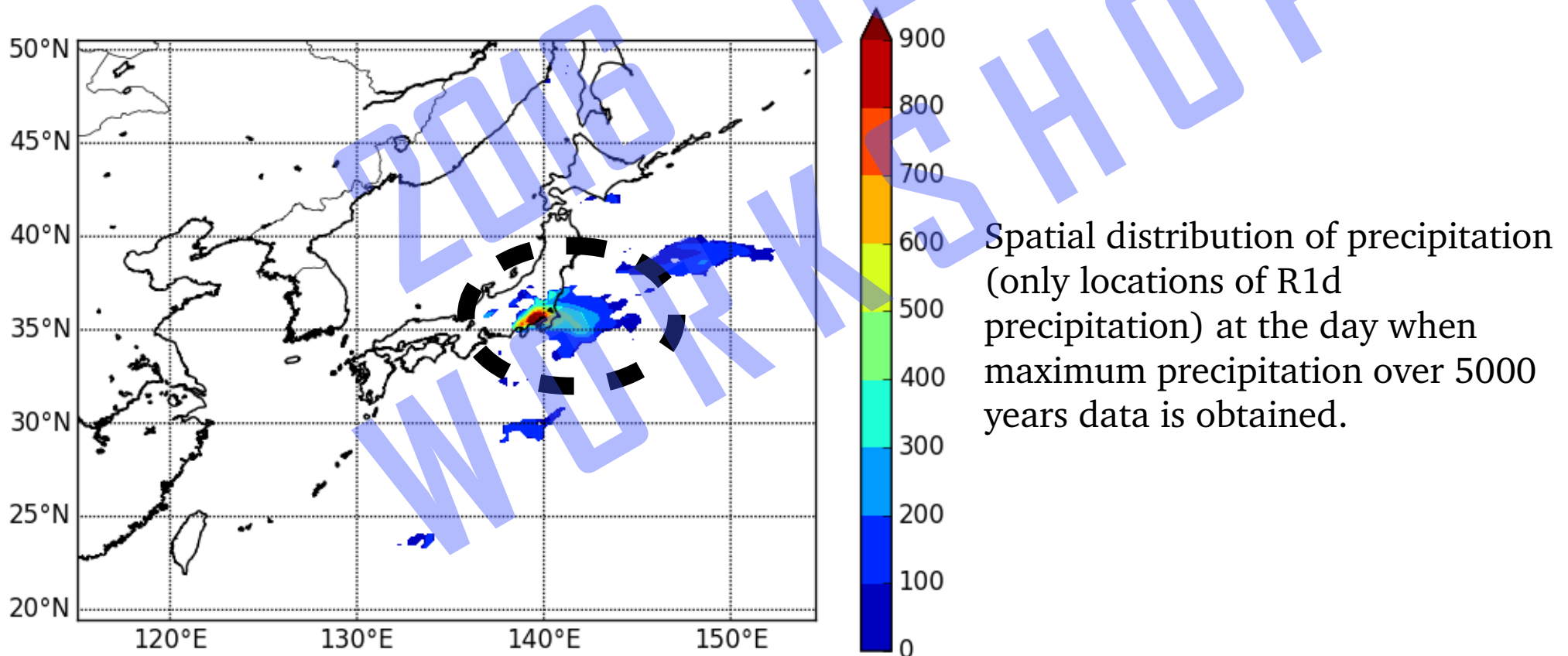
- R10d is fully followed by GEV distribution and matches the results of nonparametric method.





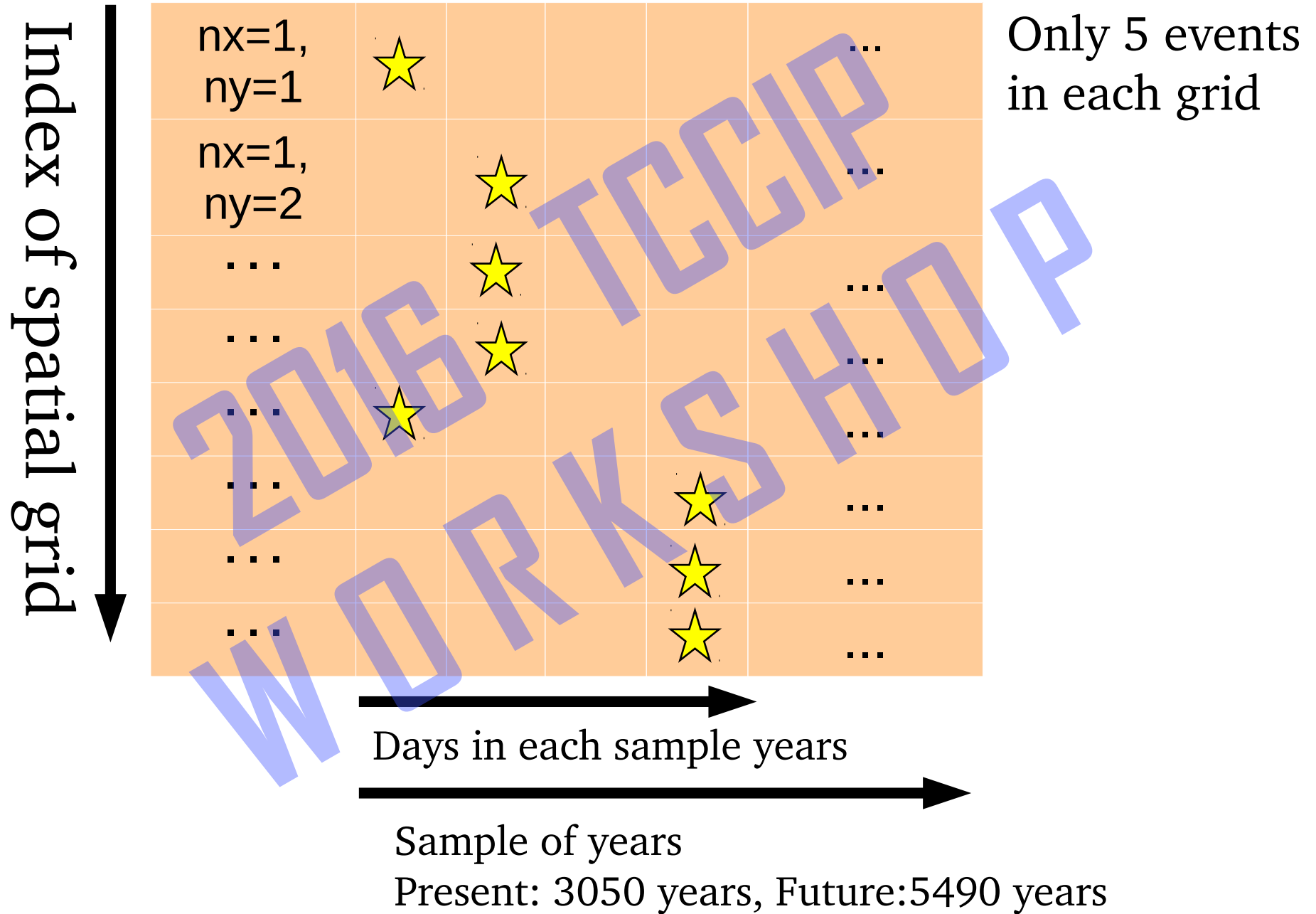
# Extreme precipitation with 1000 years return levels is a rare event?

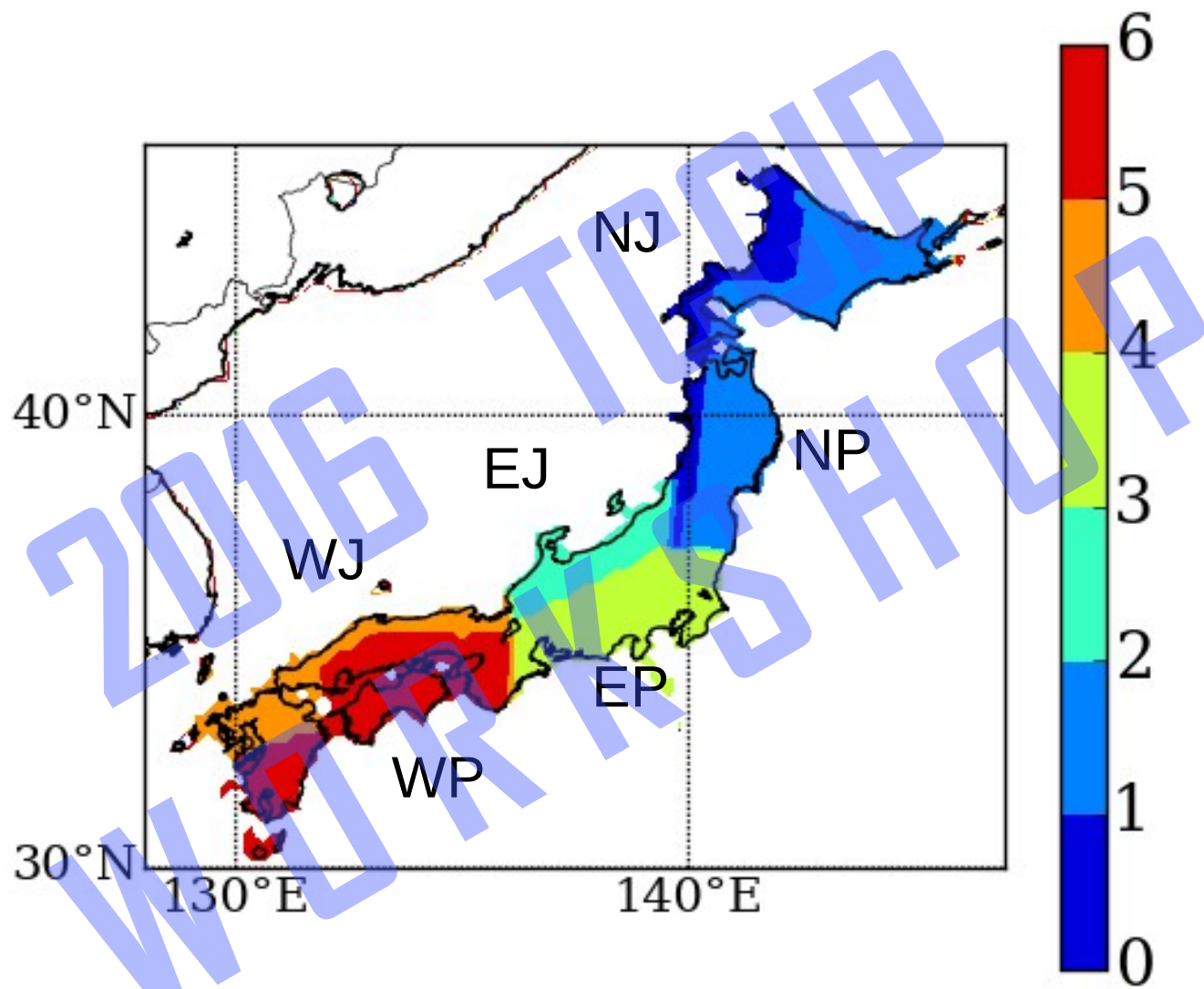
- Return levels are defined as a local estimation of extreme precipitation.
- The precipitation events usually do not extend to whole target areas (ex. Japan islands) and are limited in a small regions (e.g. bottom figure)
- Considering some areas, they happen more frequently somewhere in the areas than once for 1000 years.



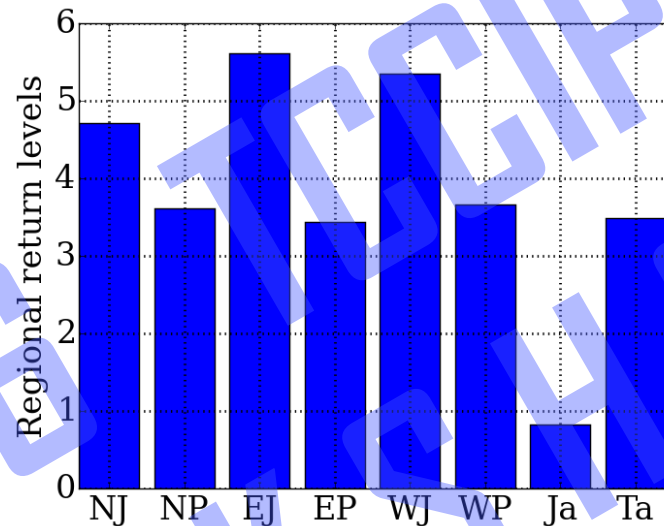
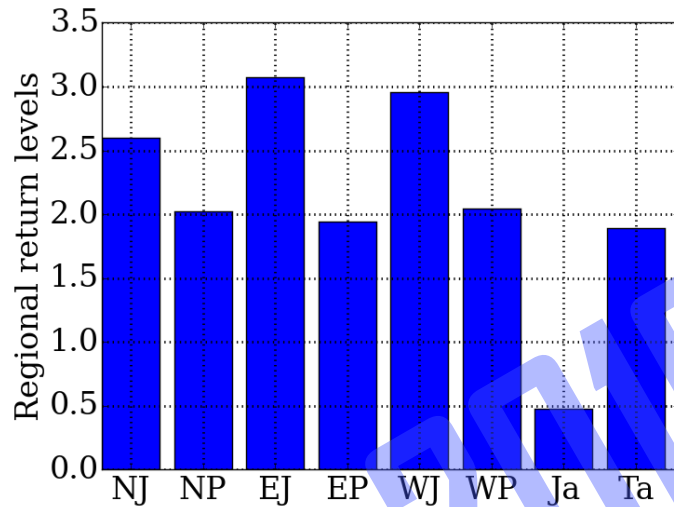
# Regional return periods

(local return levels = 1000year in future experiment)

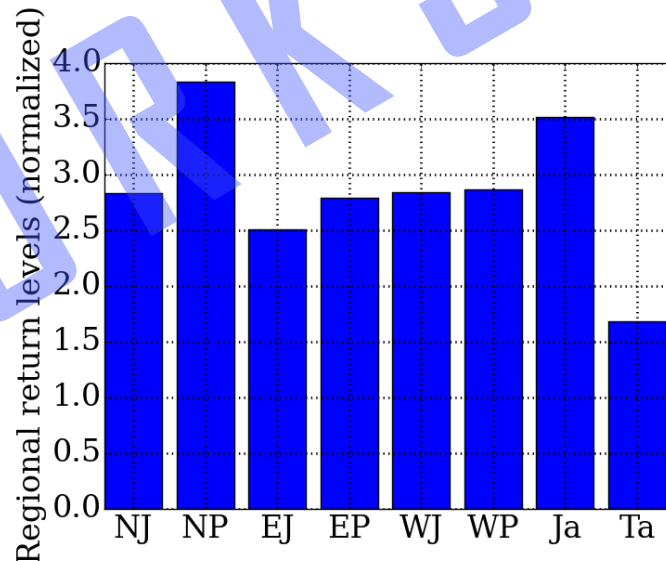
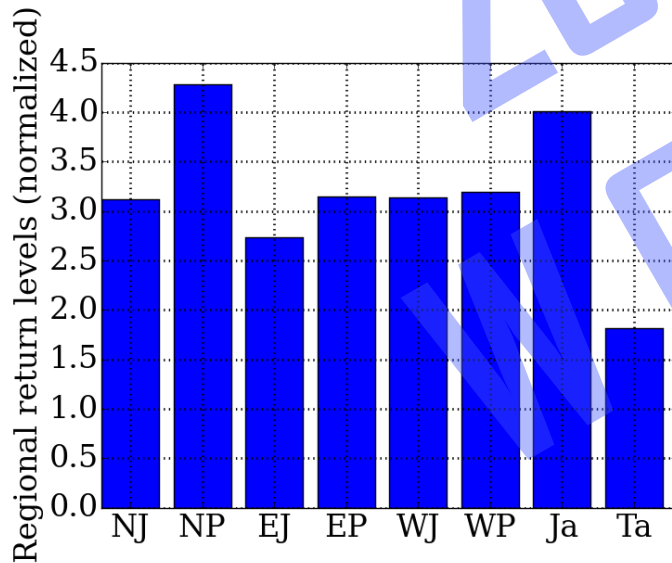




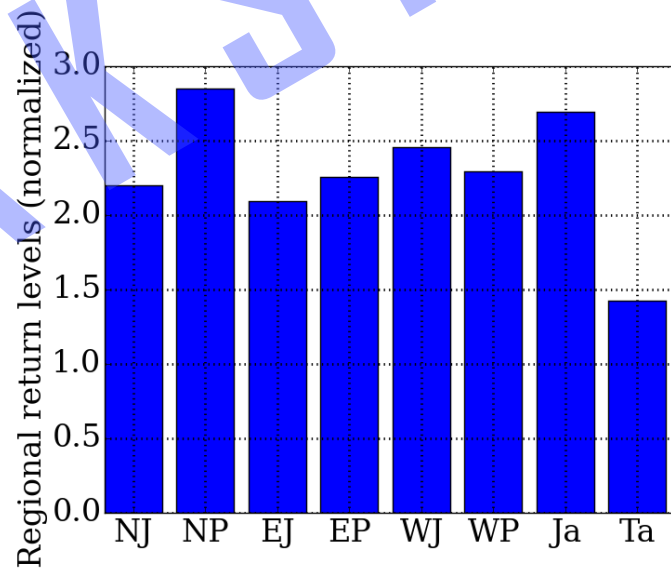
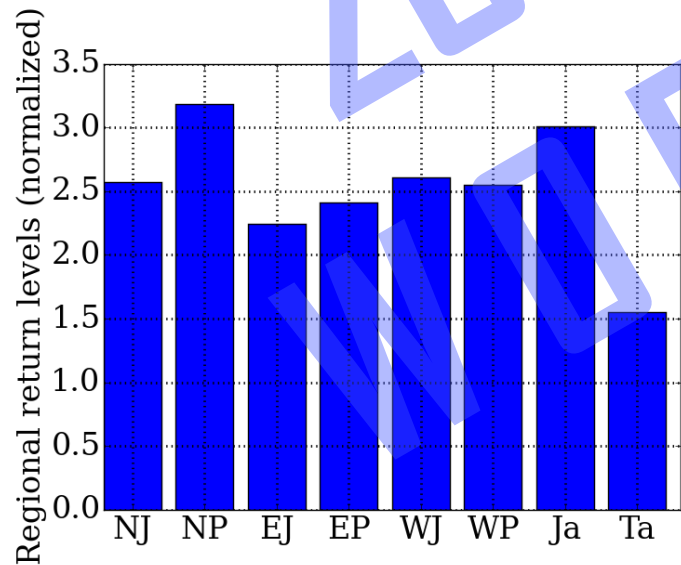
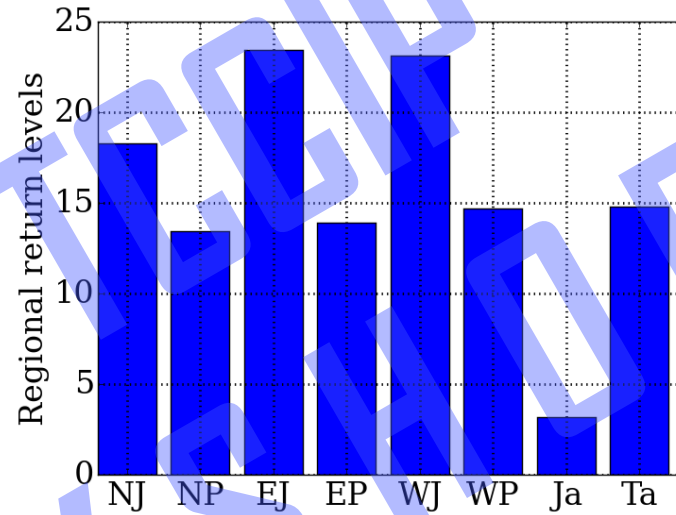
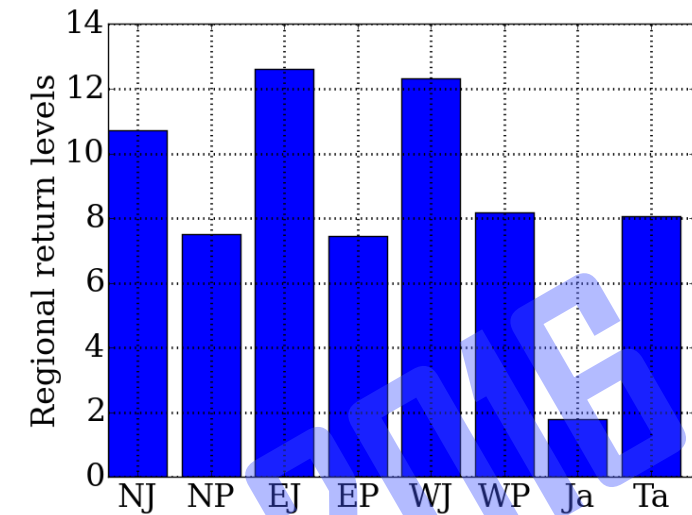
# How many days do we watch a news about extreme precipitation with X return periods?



Local return period =  
100 years (left)  
and 200 years (right)



Local return period = 500 years (left)  
and 1000 years (right)



# Summary and conclusion

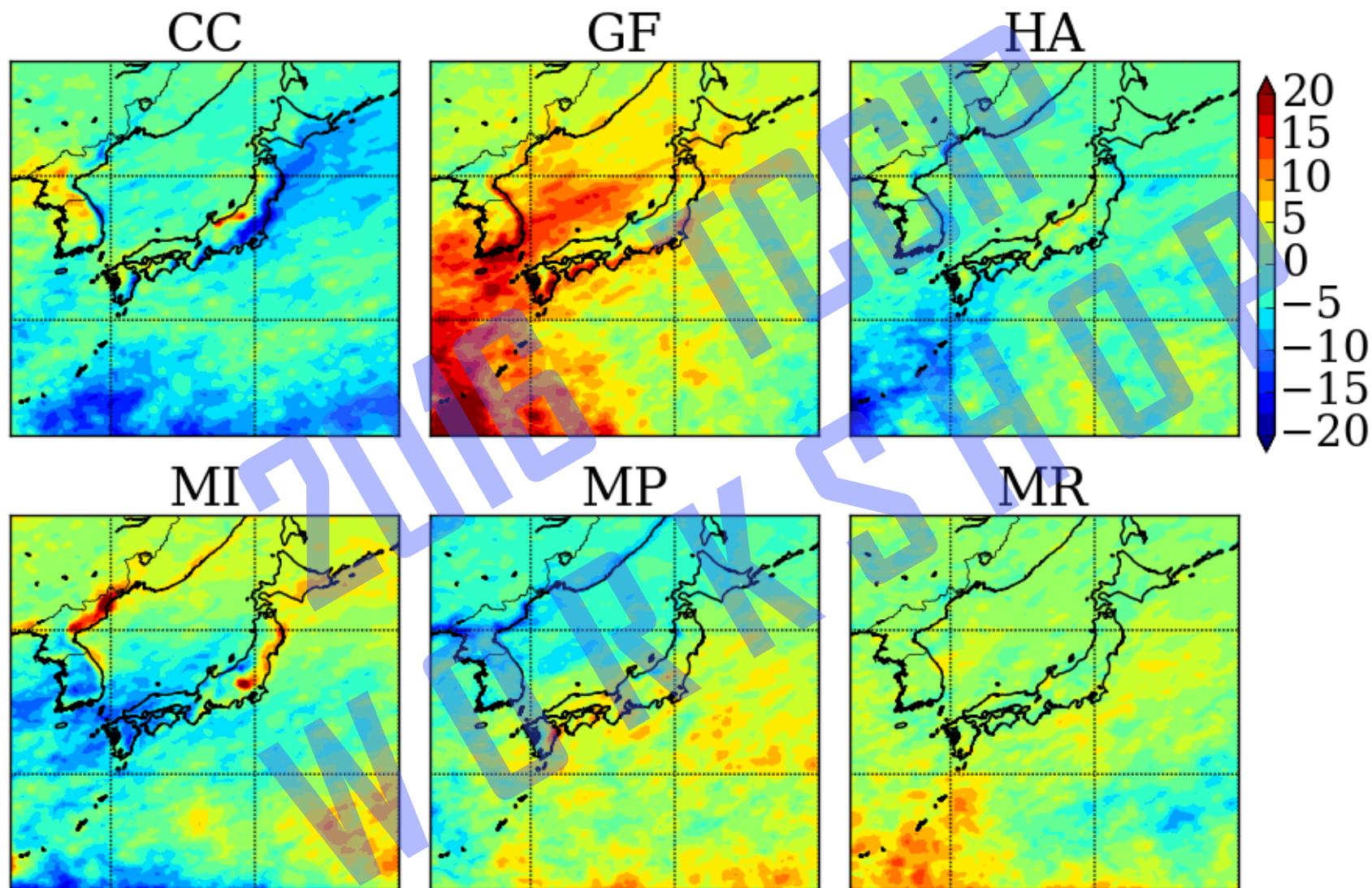
- We estimate the present and future extreme precipitation using d4PDF regional data, which enables the calculation of the very extreme precipitation of the order of 1000 years and its spatial distribution.
- Non-parametric method yields reliable estimation of extreme precipitation without any assumption about the statistical characteristics of extreme precipitation.
- The present study shows that the validity of the GEV method around Japanese islands areas is high. On the other hand, the GEV method yield overestimated values of large return levels on the ocean, which suggests the use of alternative functions to fit the R1d data (logarithmic normal, exponential etc.).
- The method using various fitting functions have technical problem about which function should be used; non-parametric method has an advantage over them.
- Non-parametric method enables the domain-accumulated frequency of extreme precipitation when given a return periods.

# Thank you for your attention

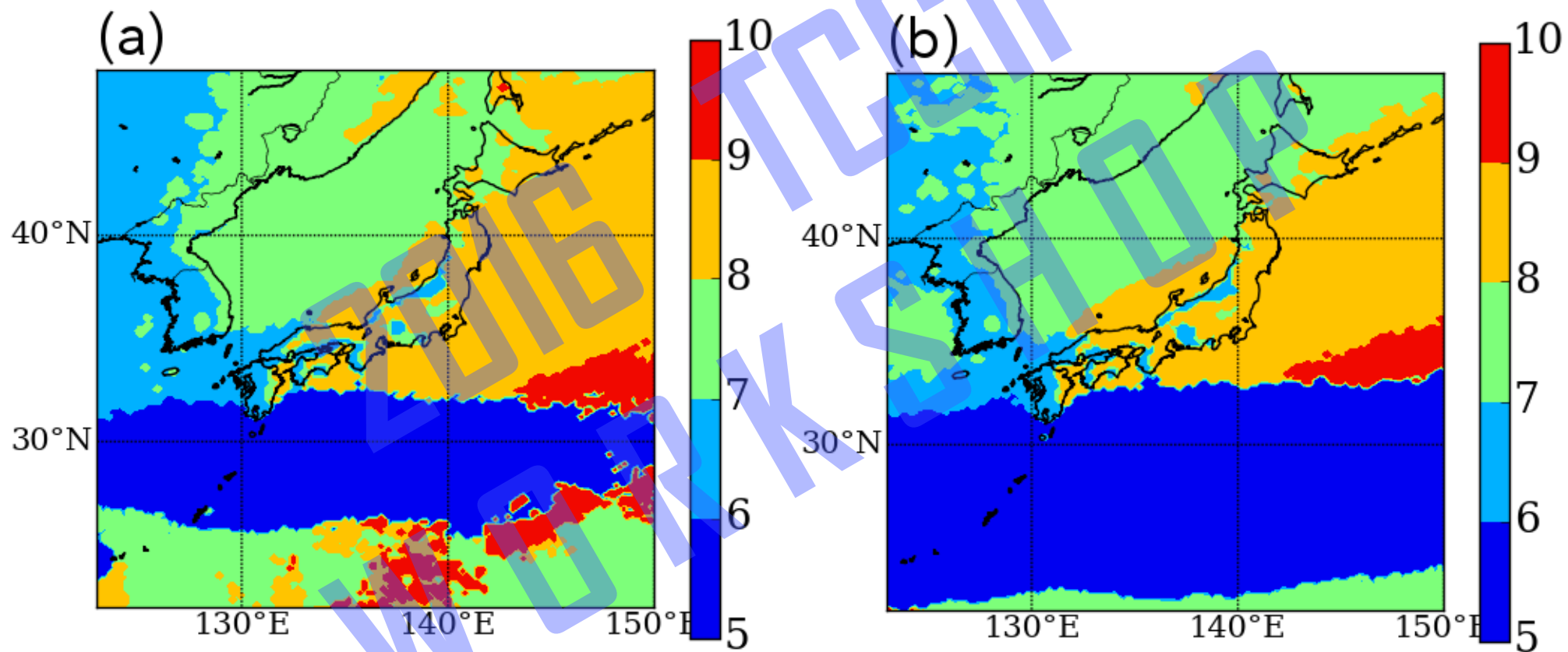
This study was supported by Ministry of Education, Culture, Sports, Science and Technology of Japan (MEXT) under the framework of the “SOUSEI” program.



# アンサンブルグループ間の違い



# R1dが最も多く取り出された月





# Introduction of “Database for Probabilistic Description of Future Climate Change (d4PDF)”

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1: Meteorological Research Institute

2: Atmosphere and Ocean Research Institute, The University of Tokyo

3: Disaster Prevention Research Institute, Kyoto University

# Climatic Hazards and Risk Assessment

What climate researcher can estimate

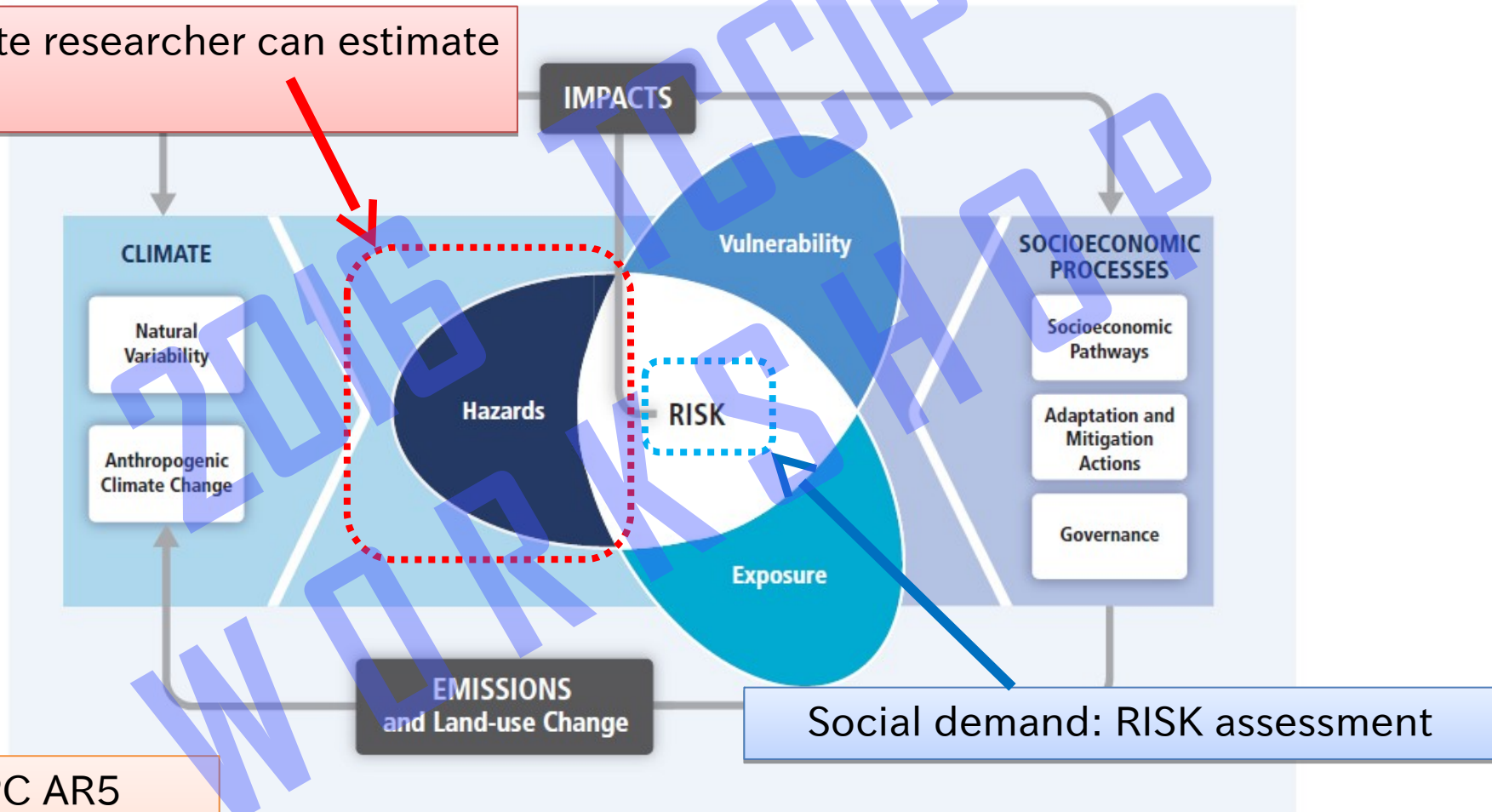


Figure of IPCC AR5 WGII

of the core concepts of the WGII AR5. Risk of climate-related impacts results from the interaction of climate-related hazards (including hazardous vulnerability and exposure of human and natural systems. Changes in both the climate system (left) and socioeconomic processes including adaptation and mitigation (right) are drivers of hazards, exposure, and vulnerability. [19.2, Figure 19-1]

Calculate many ensemble cases with a rather coarse resolution model

Calculate for specific case with a very high resolution model

## Information needed for RISK assessment

① Disaster prevention

② Disaster mitigation

Probabilistic information

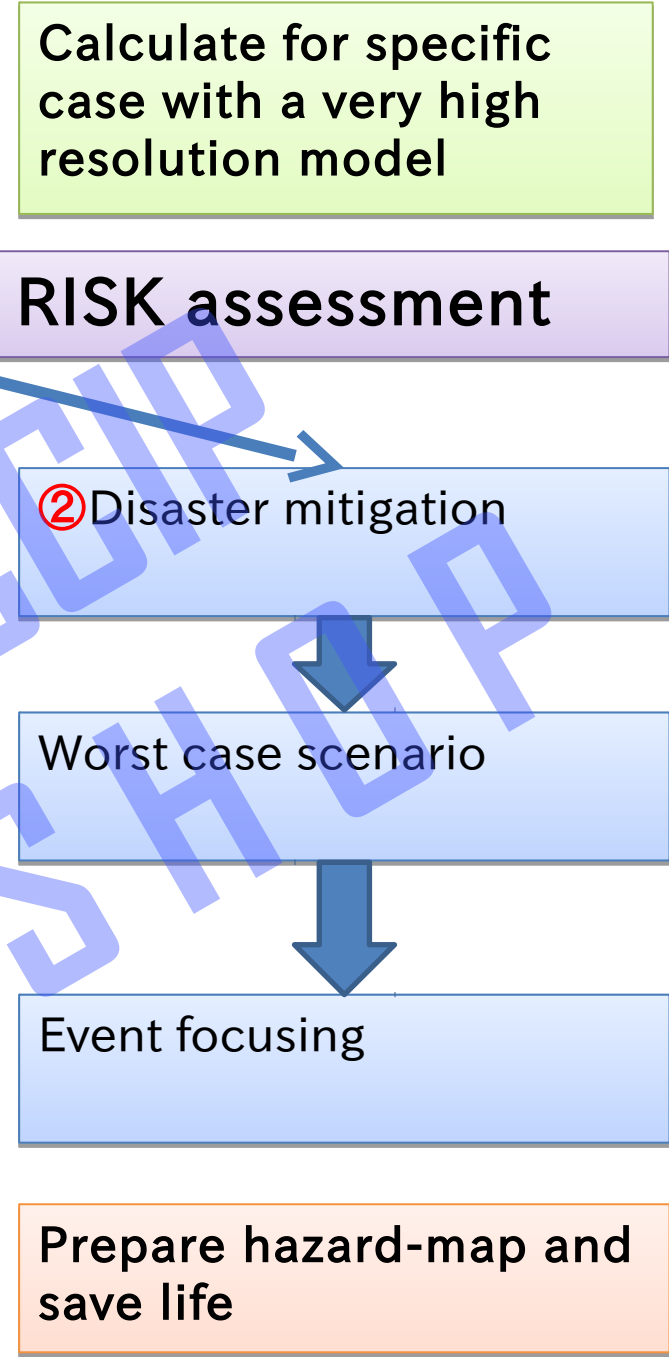
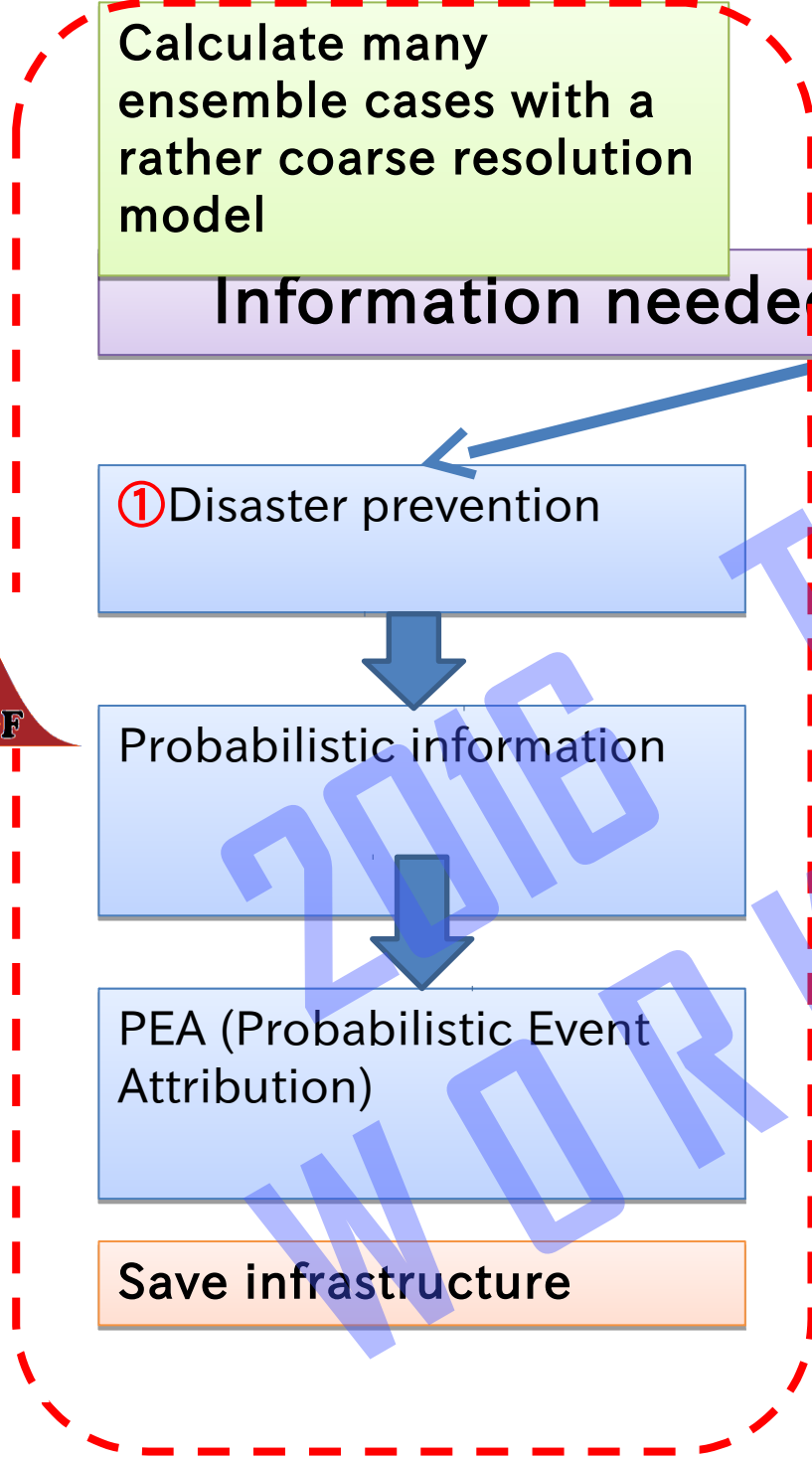
Worst case scenario

PEA (Probabilistic Event Attribution)

Event focusing

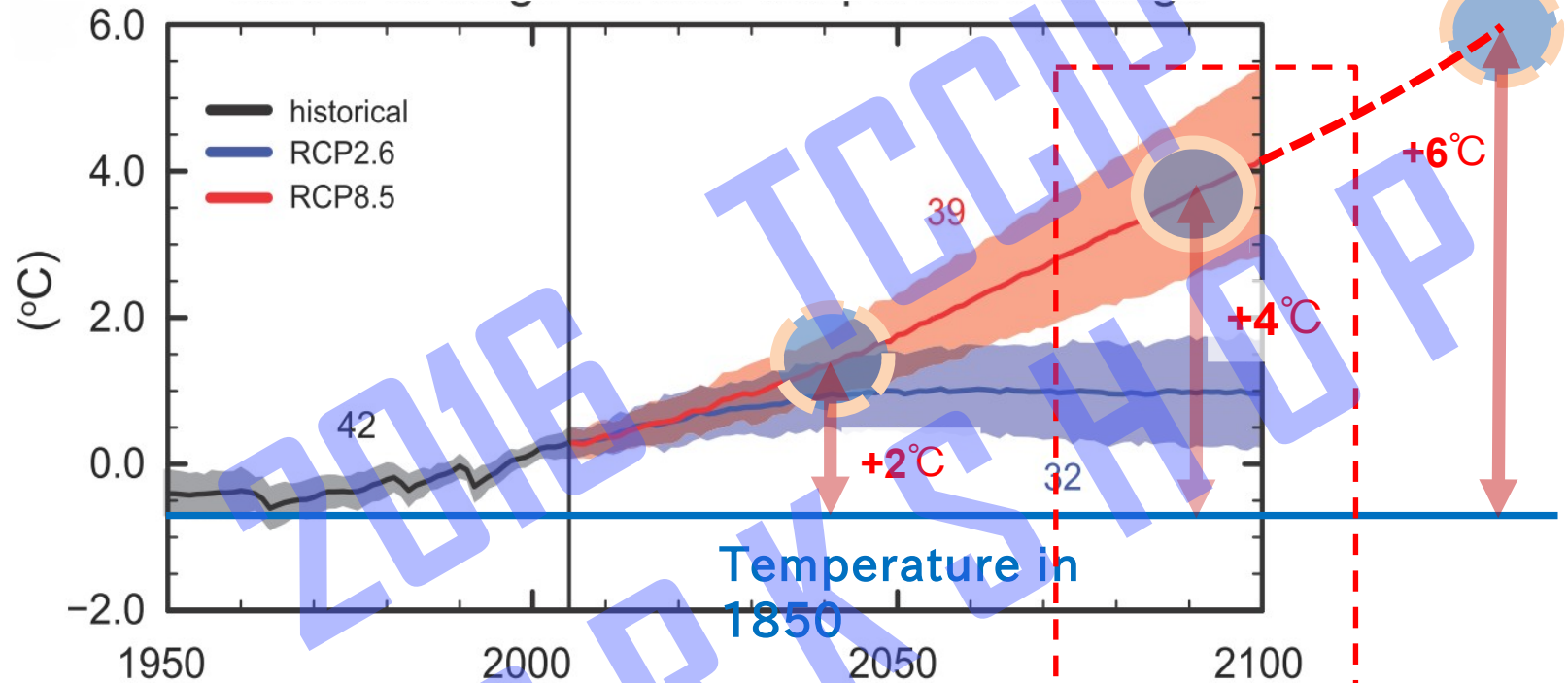
Save infrastructure

Prepare hazard-map and save life



# Design of d4PDF experiments

Surface temperature of the whole world mean



60km AGCM

100 members

90 members  
( $6\Delta T \times 15\delta T$ )

NHRCM 20km  
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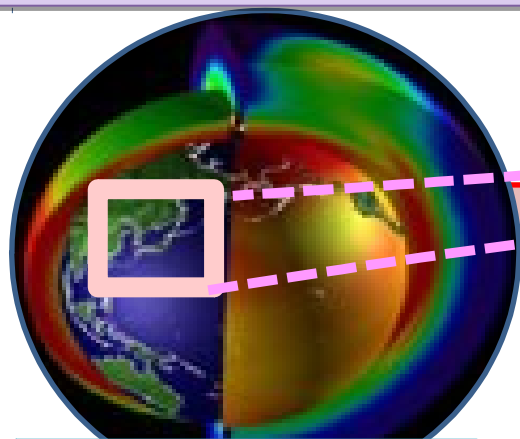
1951 Present exp. 2010

60 years

$+4^{\circ}\text{C}$   
exp.

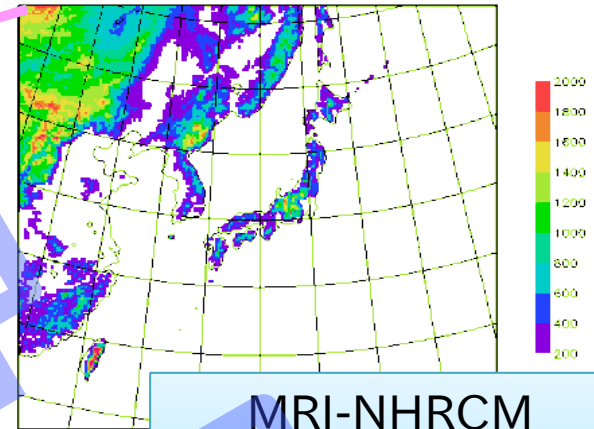


# The database which express the 5,400 patterns in +4°C world



MRI-AGCM Grid  
size: 60km

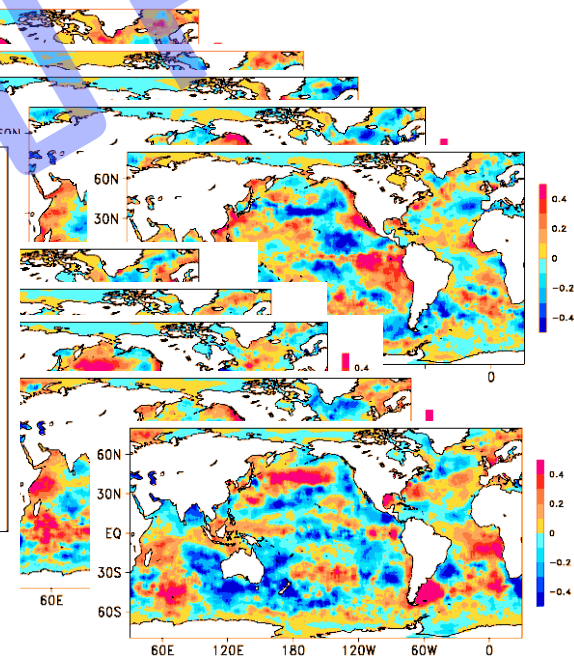
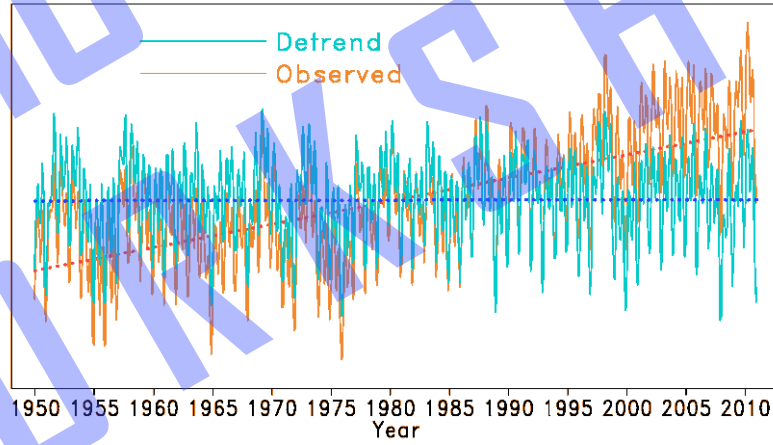
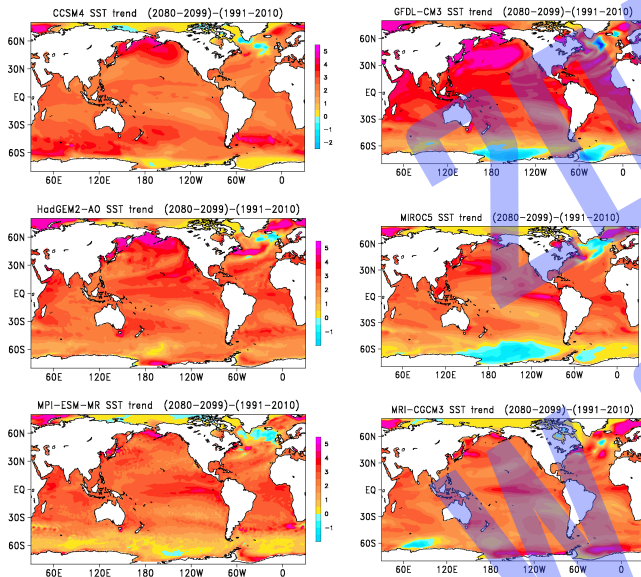
Dynamical downscaling



MRI-NHRCM  
Grid size: 20km

Data amount: 1.2 PB

## SST for +4°C experiments



6 type C.C. trends in  
CMIP5 experiments  
( $\Delta T$ ; Shiogama et al. 2010)



SST without trend  
(blue line; COBE-SST2)

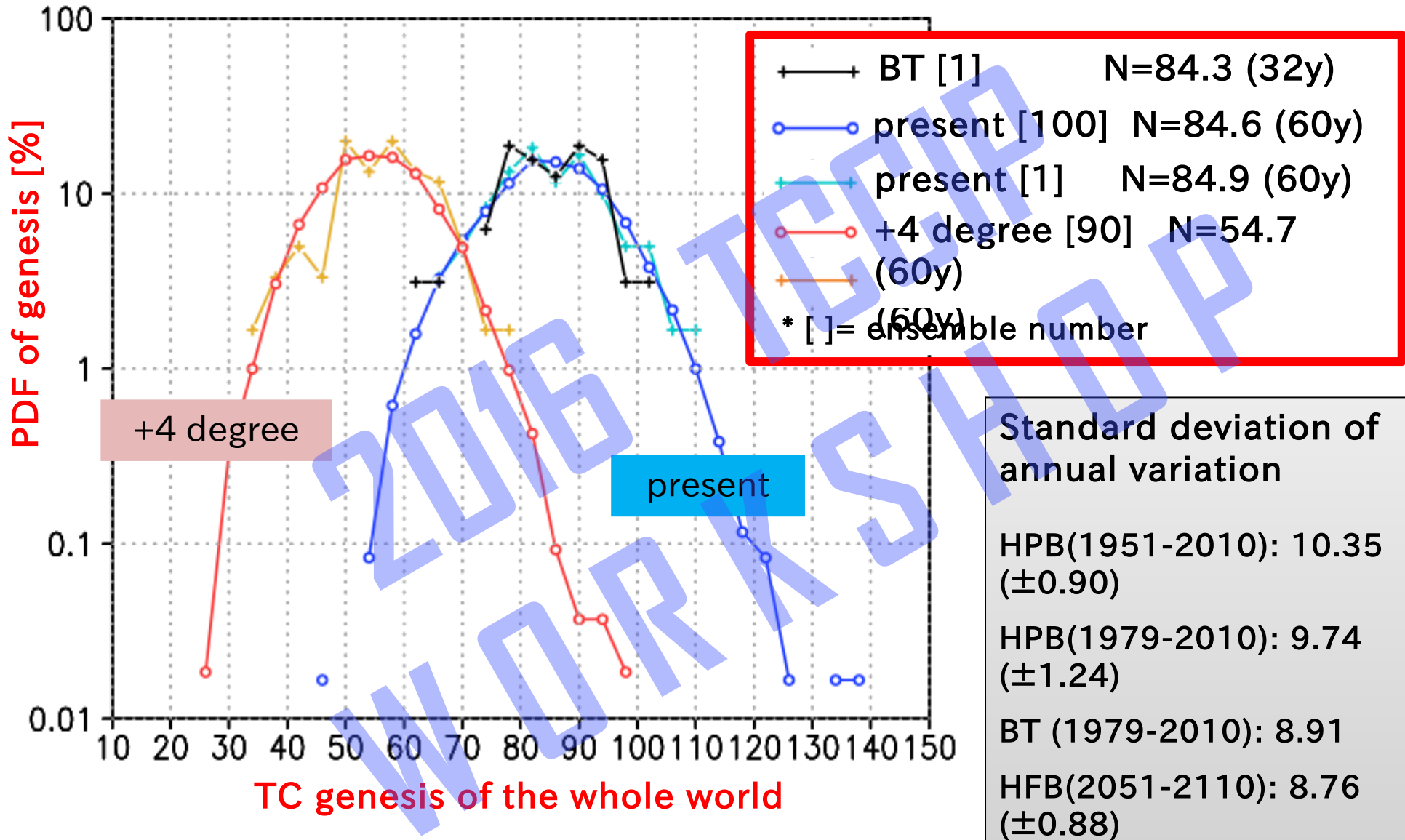


15 ensembles which  
count the uncertainty  
of observed value  
( $\delta T$ )

C.C. trends is 1<sup>st</sup> mode of EOF



# Change of TC genesis of the whole world



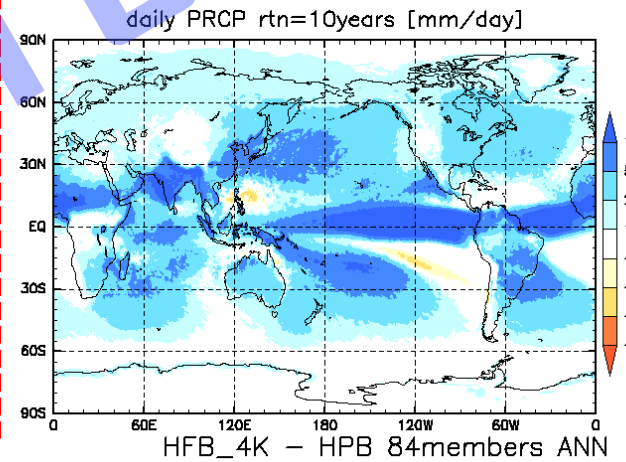
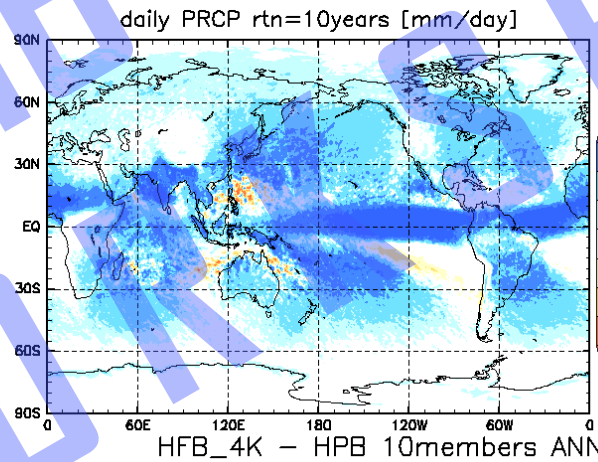
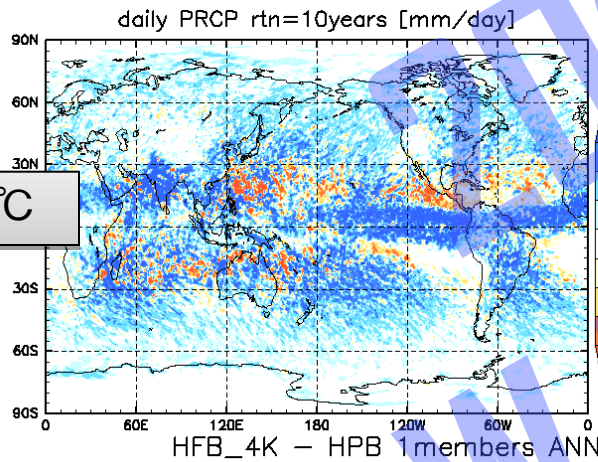
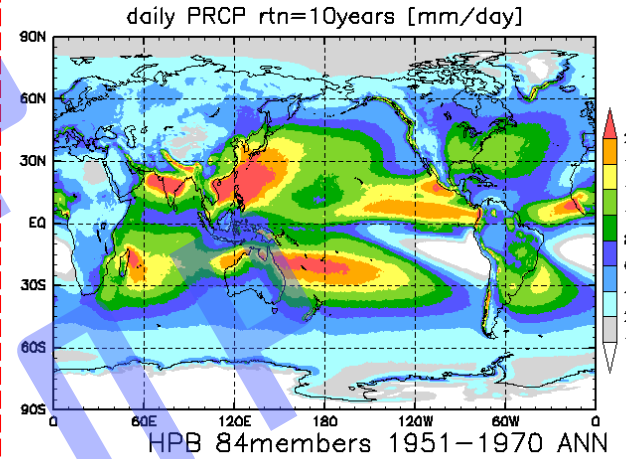
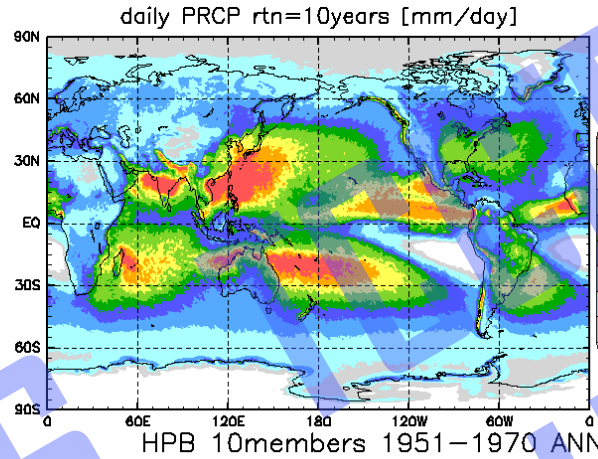
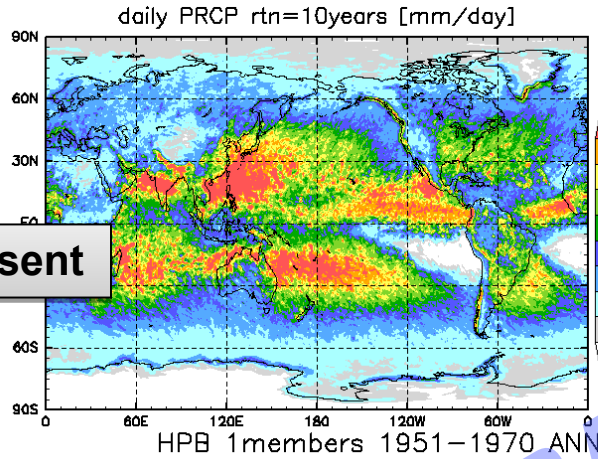
# 10 years return periods daily precipitation



1 member

10 members

84 members



present

+4°C

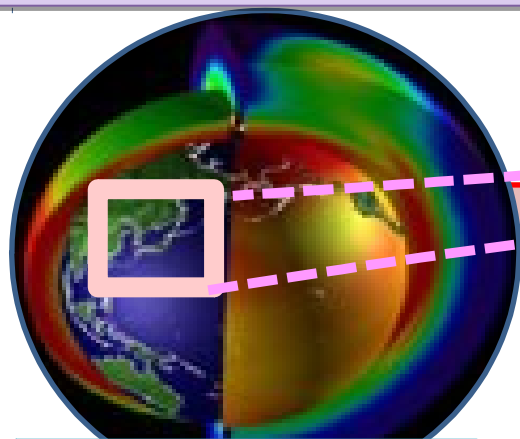
- Changes of such heavy precipitation become clear as the member increase
- Such heavy precipitation increase in the Eastern Pacific, India, and Western Africa

# Summary



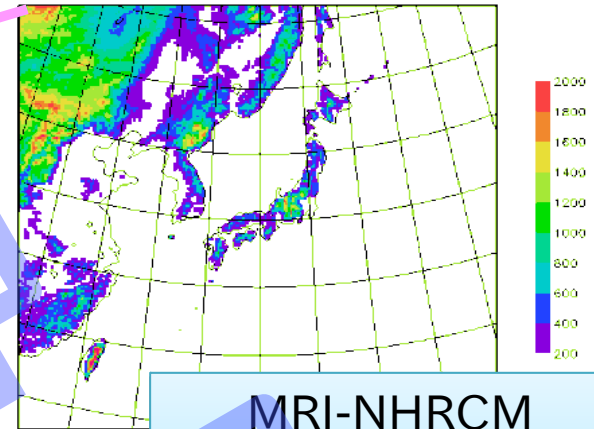
- For RISK assessment of C.C., we have to think how to calculate the C.C. by using our model.
- Here we have two schemes to get useful data for impact study researchers.
- One method is to calculate so many ensemble number to get PDF of the phenomena.
- We introduce here **d4PDF** project.

# The database which express the 5,400 patterns in +4°C world



MRI-AGCM Grid  
size: 60km

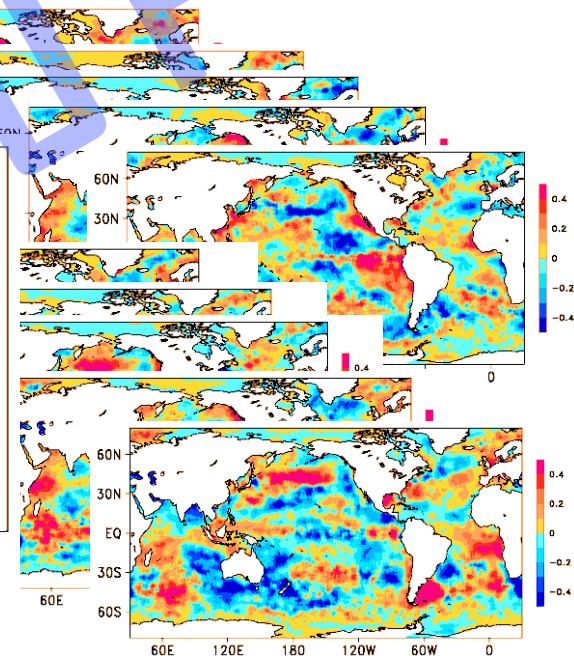
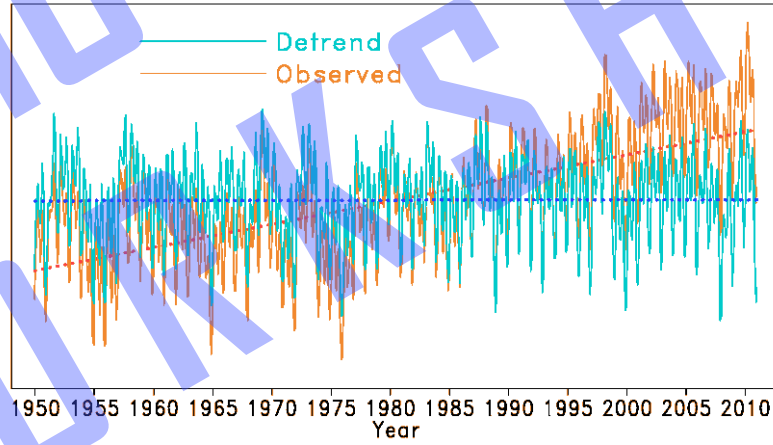
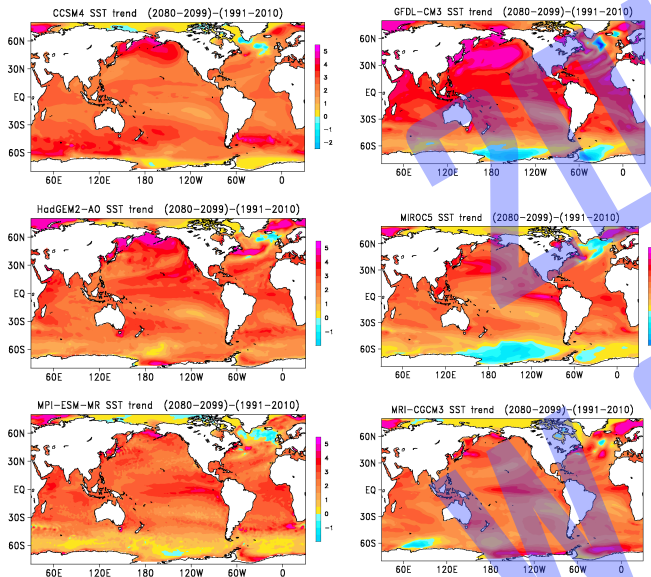
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